

# Lattice Boltzmann method for variable viscous fluid flow on spherical surface

Junxiang Yang<sup>a</sup>, Seungyeon Kang<sup>b</sup>, Youngjin Hwang<sup>b</sup>, Soobin Kwak<sup>b</sup>, Seokjun Ham<sup>b</sup>,  
Junseok Kim<sup>b,\*</sup>

<sup>a</sup> School of Computer Science and Engineering, Faculty of Innovation Engineering, Macau University of Science and Technology, Macao Special Administrative Region of China

<sup>b</sup> Department of Mathematics, Korea University, Seoul 02841, Republic of Korea

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## ABSTRACT

We propose an efficient numerical method for an incompressible fluid flow with variable viscosity on spherical surface. The proposed computational scheme is based on a finite volume lattice Boltzmann method (FVLBM). The spherical surface is triangulated and each point on the triangular mesh is assigned to one of two values of variable viscosity. Simplified coastlines using interpolation makes our proposed method highly efficient for solving fluid flows. Using the proposed algorithm, we simulate various fluid flows over inhomogeneous domains, i.e., land and sea areas. We apply different viscosity values for each domain using different relaxation times based on position of node points. Moreover, the progression of the storm is examined to demonstrate the effectiveness of the proposed approach.

## 1. Introduction

In recent years, tropical storms have occurred in many places on Earth, with some even developing into hurricanes, cyclones, and typhoons. These natural disasters cause a catastrophic amount of damage, particularly in coastal regions. Accordingly, there has been an urgent quest for fast and accurate storm event models. In this paper, we present a new method for analyzing storm behavior on the Earth's surface. Our method solves the incompressible surface Navier–Stokes (NS) equation on surface  $S$  using a finite volume lattice Boltzmann method (LBM).

LBM is a highly effective tool in computational fluid dynamics (CFD) technology. An earlier review of the application of LBM in fluid flow can be tracked back to [1]. Later, Aidun and Clausen [2] introduced developments of LBM on complex fluids before the year 2010. The LBM was developed based on an earlier idea of Lattice-Gas Cellular Automata; Wolf-Gladrow [3] reviewed these two methods in detail. Similarly, Succi [4] also reviewed the fluid models of LBM and introduced some potential applications. For the theories and engineering practices of LBM, interested readers can refer to [5,6]. Huang et al. [7] introduced practical LBM in multi-phase fluid simulations. As for a recent review of phase-field LBM in multi-phase fluid applications, please refer to [8] and references therein.

In recent years, some new progresses of LBM appears in fluid flow problems and aerodynamics. The phase-field based LBM [8–10] for

multiphase flows were studied. The study of the mixed convection of nanofluids flow [11–13], garnered significant attention. Researchers also investigated laminar flow in both two-dimensional (2D) [13] and three-dimensional (3D) [14] domain. Furthermore, Mandzhieva and Subhankulova [15] applied the lattice Boltzmann model to study the pore-scale models. Babanezhad et al. [16] simulated the 3D cavity flow of a large dataset for the build-up fluid flow using LBM. Li et al. [17] introduced a LBM to describe interaction of waves and porous structures in 3D space. Hosseini et al. [18] reviewed various entropic LBMs and presented their recent progress and important challenges in that area.

The conventional LBM basically consists of two steps: moving the particle and relaxing distribution functions towards equilibrium. This coupling results restricts the LBM to Cartesian meshes [19–21]. Several extensions of LBM have been developed to overcome such restrictions. One of the pioneering work was done by He et al. [22], who added an interpolation step to the conventional LBM. Another noteworthy study is by Nannelli et al. [23], where Nannelli proposed a finite volume scheme for LBM. Since then, many works have been done using the finite volume scheme for LBM [24–26].

The above-mentioned studies only dealt with a fixed viscosity value. When substituting Earth into the problem, viscosity on land and sea has to be different. Therefore, we should solve LBM on a surface with various viscosity values. Here are some studies that considered surfaces

\* Corresponding author.

E-mail address: [cfdkim@korea.ac.kr](mailto:cfdkim@korea.ac.kr) (J. Kim).

URL: <https://mathematicians.korea.ac.kr/cfdkim/> (J. Kim).

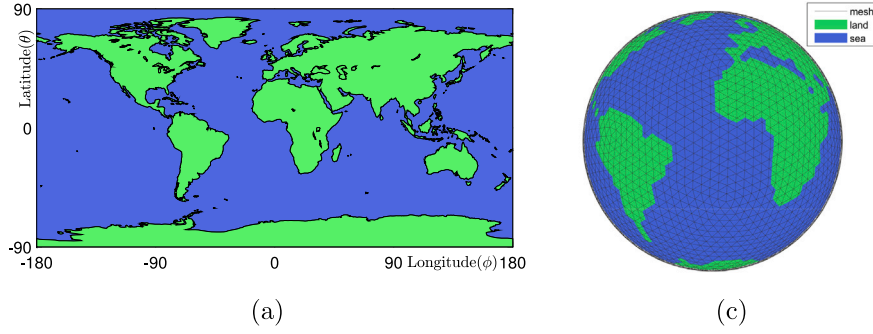


Fig. 1. (a) The world map in 2D space and (b) mapping on a spherical surface mesh. Here, green and blue colors denote land and sea regions, respectively.

with multiple viscosity values. Hussain et al. [27] demonstrated the flow behavior of a Casson fluid with variable viscosity that exhibits steady-state flow in 2D. The system is converted to an ODE, and numerical solutions are proposed. Chen et al. [28] presented an immersed boundary-simplified LBM for simulations of incompressible viscous flows. In this study, the relation between the relaxation parameter  $\tau$  and viscosity  $\nu$  is given as  $\nu = c_s^2(\tau - \frac{1}{2})\delta_t$ . Guan and Novoselov [29] studied a numerical investigation of electroconvection phenomena between two parallel plates and proposed a two-relaxation time. The LBM is applied to fluids in porous media [30,31]. He et al. [30] studied the heat and mass transfer characteristics of a fluid with temperature-dependent viscosity using LBM. Li and Dimitrienko [31] studied the generalized Newtonian fluid in porous media using the homogenization method. After developing numerical methods, Li and Dimitrienko demonstrated the effects of viscosity on local problems. There are application of LBM to Navier–Stokes and convection–diffusion equations [32,33]. Chai and Baochang [32] presented a unified framework of multiple-relaxation-time LBM. Their model is depicted with a matrix that allows decoupled relaxation. Zhao et al. [33] proposed a block triple-relaxation-time LBM. One of the striking features of this model is that the model is partitioned into three relaxation parameter blocks. The relation between relaxation parameters is determined based on the analysis of the half-way bounce-back scheme.

In the previous work, we developed a finite volume scheme for the LBM on curved surfaces in 3D space [34], which used a constant viscosity. The main aim of this paper is to suggest an efficient numerical method for an incompressible fluid flow with variable viscosity on spherical surface so that we can simulate real-world scenarios where fluid viscosity changes with temperature or other fields.

The paper is organized as follows. In Section 2, we introduce the discretization of coastlines, which are simplified for calculation purpose. In Section 3, we present numerical results using LBM. Finally, in Section 4, we provide a brief conclusion.

## 2. Numerical method

### 2.1. Discretization of the computational domain

Fig. 1(a) and (b) illustrate the world map in 2D space and mapping on a spherical surface mesh, respectively. Green and blue colors denote land and sea regions, respectively. We use the ‘load coastlines’ command in the Mapping Toolbox to create a world map in MATLAB [35]. The ‘load coastlines’ command returns the coastline data in latitude (‘coastlat’) and longitude (‘coastlon’). See Fig. 2(a) for the plot of coastlines, where the horizontal and vertical axes represent longitude  $\phi$  and latitude  $\theta$ , respectively.

For simplicity, we exclude small islands by removing coastlines with fewer than 100 points and coastlines that lie within another coastline. As shown in Fig. 2(b), this process results in only nine large coastlines. We further simplify the coastlines using an interpolation method and reduce the number of points for each coastline by a factor

of five. Smooth interpolated coastlines are shown in Fig. 2(c). The spherical surface is discretized using a triangular mesh, as illustrated in Fig. 2(d). At each vertex point  $(x, y, z)$  of the triangular mesh, we find the corresponding angles  $(\theta, \phi) = (\arcsin(z/R), \arctan(y/x))$ , where  $\theta$ ,  $\phi$ , and  $R$  are schematically illustrated in Fig. 2(e). Next, we find whether  $(\theta, \phi)$  is in a land or sea area. We define an indicator function  $\mathbb{I}_c$  such that if  $(x, y, z)$  is in land,  $\mathbb{I}_c(x, y, z) = 1$  and  $\mathbb{I}_c(x, y, z) = 0$  when  $(x, y, z)$  is in the sea. Figs. 2(f) and (g) illustrate  $(x, y, z)$  points in green when  $\mathbb{I}_c(x, y, z) = 1$  and in blue when  $\mathbb{I}_c(x, y, z) = 0$ , on 2D space and the spherical surface, respectively.

### 2.2. Finite volume lattice Boltzmann method

In this paper, we simulate variable viscous and incompressible fluid flow on a spherical surface by modifying the finite volume lattice Boltzmann equation proposed by Yang et al. [34]. Among the various methods for solving the lattice Boltzmann equation, we choose the finite volume method due to its applicability to arbitrarily shaped domains and irregular meshes [36]. In particular, we apply this method to a curved spherical surface. A disadvantage is its increased complex compared to the conventional LBM.

The governing equation of finite volume lattice Boltzmann method (FVLBM) is given as follows [37]:

$$\frac{\partial f_i}{\partial t}(\mathbf{x}, t) + \mathbf{c}_i \cdot \nabla f_i(\mathbf{x}, t) = -\frac{1}{\tau(\mathbf{x})} \left( f_i(\mathbf{x}, t) - f_i^{eq}(\mathbf{x}, t) \right), \quad i = 1, 2, \dots, 19. \quad (1)$$

For  $\mathbf{x} = (x, y, z)$  and time  $t$ ,  $f_i(\mathbf{x}, t)$  and  $f_i^{eq}(\mathbf{x}, t)$  are particle distribution function and local equilibrium distribution, respectively.  $i$ th directional velocity is  $\mathbf{c}_i$ . For  $f_i(\mathbf{x}, t)$  and  $\mathbf{c}_i$ , index  $i$  varies from 1 to 19.  $\tau(\mathbf{x})$  is the relaxation time at space  $\mathbf{x}$ . Values of  $\mathbf{c}_i$  are given as follows:

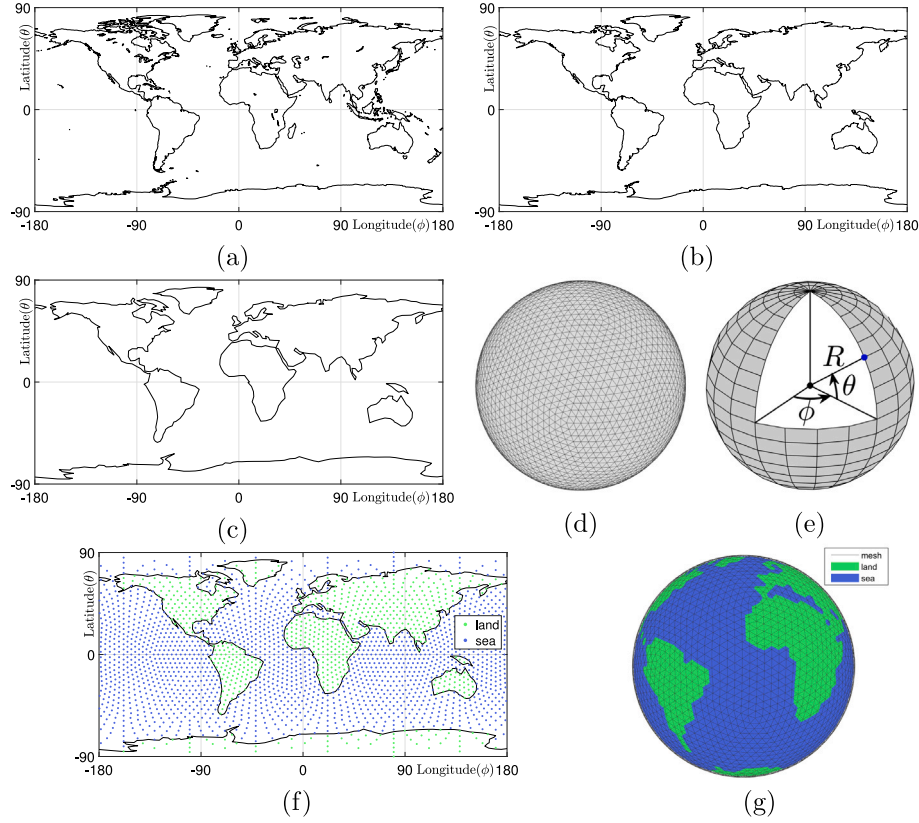
$$\begin{aligned} \mathbf{c}_1 &= (1, 0, 0), \mathbf{c}_2 = (-1, 0, 0), \mathbf{c}_3 = (0, 1, 0), \mathbf{c}_4 = (0, -1, 0), \\ \mathbf{c}_5 &= (0, 0, 1), \mathbf{c}_6 = (0, 0, -1), \mathbf{c}_7 = (1, 1, 0), \mathbf{c}_8 = (1, -1, 0), \\ \mathbf{c}_9 &= (1, 0, 1), \mathbf{c}_{10} = (1, 0, -1), \mathbf{c}_{11} = (-1, 1, 0), \mathbf{c}_{12} = (-1, -1, 0), \\ \mathbf{c}_{13} &= (-1, 0, 1), \mathbf{c}_{14} = (-1, 0, -1), \mathbf{c}_{15} = (0, 1, 1), \mathbf{c}_{16} = (0, 1, -1), \\ \mathbf{c}_{17} &= (0, -1, 1), \mathbf{c}_{18} = (0, -1, -1), \mathbf{c}_{19} = (0, 0, 0). \end{aligned}$$

Eq. (1) is solved on surface  $S$  by discretizing the surface into a triangular mesh, shown in Fig. 3(a). Consider a point  $\mathbf{P}$  on the mesh and points  $\mathbf{P}_i$ ,  $i = 1, 2, \dots$ , which are one-ring neighboring points of  $\mathbf{P}$ . In case of Fig. 3(a),  $i$  ranges from 1 to 6. Let  $\mathbf{M}$  be the center of edge  $\mathbf{P}\mathbf{P}_1$  and  $\mathbf{G}$  be the centroid of triangle  $\triangle \mathbf{P}\mathbf{P}_1\mathbf{P}_2$ . Integrating Eq. (1) over triangle  $\triangle \mathbf{P}\mathbf{M}\mathbf{G}$ , we obtain the following equation:

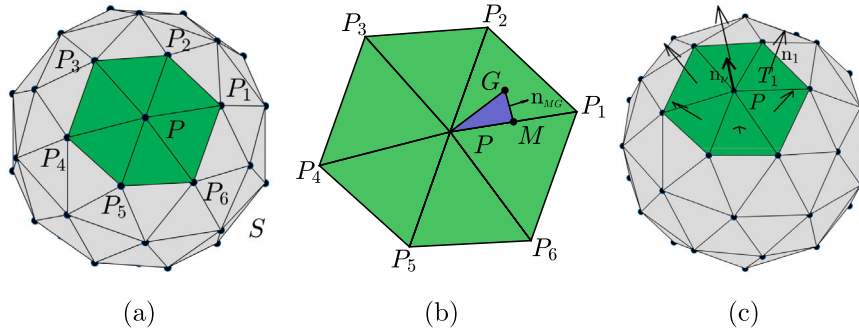
$$\begin{aligned} \int_{\triangle \mathbf{P}\mathbf{M}\mathbf{G}} \frac{\partial f_i}{\partial t}(\mathbf{x}, t) d\sigma + \int_{\triangle \mathbf{P}\mathbf{M}\mathbf{G}} \mathbf{c}_i \cdot \nabla f_i(\mathbf{x}, t) d\sigma \\ = - \int_{\triangle \mathbf{P}\mathbf{M}\mathbf{G}} \frac{1}{\tau(\mathbf{x})} \left( f_i(\mathbf{x}, t) - f_i^{eq}(\mathbf{x}, t) \right) d\sigma. \end{aligned} \quad (2)$$

We approximate and calculate each term of Eq. (2). The first term of left hand side can be approximated using  $\frac{\partial f_i}{\partial t}(\mathbf{x}, t) \approx \frac{\partial f_i}{\partial t}(\mathbf{P}, t)$  for  $\mathbf{x} \in \triangle \mathbf{P}\mathbf{M}\mathbf{G}$ .  $A_{\triangle \mathbf{P}\mathbf{M}\mathbf{G}}$  is the area of triangle  $\triangle \mathbf{P}\mathbf{M}\mathbf{G}$  at time  $t$ .

$$\int_{\triangle \mathbf{P}\mathbf{M}\mathbf{G}} \frac{\partial f_i}{\partial t}(\mathbf{x}, t) d\sigma \approx \int_{\triangle \mathbf{P}\mathbf{M}\mathbf{G}} \frac{\partial f_i}{\partial t}(\mathbf{P}, t) d\sigma = \frac{\partial f_i}{\partial t}(\mathbf{P}, t) \int_{\triangle \mathbf{P}\mathbf{M}\mathbf{G}} d\sigma \quad (3)$$



**Fig. 2.** (a) World Map using Mapping Toolbox. (b) Excluding small islands. (c) Applying an interpolation method for simplicity. (d) Discretizing the spherical surface using a triangular mesh. (e) Schematic illustration of  $\theta$  and  $\phi$ . (f) Illustration of  $(x, y, z)$  points in 2D space. (g) Illustration of  $(x, y, z)$  points on a spherical surface.



**Fig. 3.** Schematic diagram of (a) Triangular mesh. (b) Integration over triangle  $\Delta PMG$ . (c) velocity field correction.

$$= \frac{\partial f_i}{\partial t}(\mathbf{P}, t) A_{\Delta PMG} = \frac{f_i(\mathbf{P}, t + \Delta t) - f_i(\mathbf{P}, t)}{\Delta t} A_{\Delta PMG}.$$

The second term of left hand side of Eq. (2) is the flux, which can be calculated using the divergence theorem.  $\mathbf{n}_{PM}$ ,  $\mathbf{n}_{MG}$  and  $\mathbf{n}_{GP}$  are outward normal vectors of edge  $PM$ ,  $MG$ , and  $GP$ , respectively. See Fig. 3(b) for schematic diagram of  $\mathbf{n}$ .  $l_{PM}$ ,  $l_{MG}$ , and  $l_{GP}$  are length of  $PM$ ,  $MG$ , and  $GP$ , respectively.

$$\begin{aligned} \int_{\Delta PMG} \mathbf{c}_i \cdot \nabla f_i(\mathbf{x}, t) d\sigma &= \int_{\Delta PMG} \nabla \cdot [f_i(\mathbf{x}, t) \mathbf{c}_i] d\sigma = \int_{\partial \Delta PMG} f_i(\mathbf{x}, t) \mathbf{c}_i \cdot \mathbf{n} dl \\ &= \int_{PM} f_i(\mathbf{x}, t) \mathbf{c}_i \cdot \mathbf{n}_{PM} dl + \int_{MG} f_i(\mathbf{x}, t) \mathbf{c}_i \cdot \mathbf{n}_{MG} dl + \int_{GP} f_i(\mathbf{x}, t) \mathbf{c}_i \cdot \mathbf{n}_{GP} dl \\ &= \frac{f_i(\mathbf{P}, t) + f_i(\mathbf{M}, t)}{2} \mathbf{c}_i \cdot \mathbf{n}_{PM} l_{PM} + \frac{f_i(\mathbf{M}, t) + f_i(\mathbf{G}, t)}{2} \mathbf{c}_i \cdot \mathbf{n}_{MG} l_{MG} \quad (4) \\ &+ \frac{f_i(\mathbf{G}, t) + f_i(\mathbf{P}, t)}{2} \mathbf{c}_i \cdot \mathbf{n}_{GP} l_{GP}. \end{aligned}$$

$\mathbf{n}_{MG}$  can be calculated using edges of triangle  $\Delta PMG$ . Other outward normal vectors can be calculated in the same way.

$$\mathbf{n}_{MG} = \frac{\overrightarrow{PG} - \overrightarrow{PG} \cdot \overrightarrow{MG} \frac{\overrightarrow{MG}}{|\overrightarrow{MG}|}}{\left| \overrightarrow{PG} - \overrightarrow{PG} \cdot \overrightarrow{MG} \frac{\overrightarrow{MG}}{|\overrightarrow{MG}|} \right|}.$$

Due to the fact that  $M$  is the center of  $PP_1$  and  $G$  is centroid of triangle  $\Delta PMG$ , values of  $f_i(\mathbf{M}, t)$  and  $f_i(\mathbf{G}, t)$  can be calculated using interpolation.

$$f_i(\mathbf{M}, t) = \frac{f_i(\mathbf{P}, t) + f_i(\mathbf{P}_1, t)}{2}, \quad (5)$$

$$f_i(\mathbf{G}, t) = \frac{f_i(\mathbf{P}, t) + f_i(\mathbf{P}_1, t) + f_i(\mathbf{P}_2, t)}{3}. \quad (6)$$

On right hand side of Eq. (2), collision is approximated using interpolation similar to Eq. (6).

$$-\int_{\Delta PMG} \frac{1}{\tau(\mathbf{x})} (f_i(\mathbf{x}, t) - f_i^{eq}(\mathbf{x}, t)) d\sigma \\ = -\frac{A_{\Delta PMG}}{\tau(\mathbf{P})} \left( \frac{f_i(\mathbf{P}, t) + f_i(\mathbf{M}, t) + f_i(\mathbf{G}, t)}{3} - \frac{f_i^{eq}(\mathbf{P}, t) + f_i^{eq}(\mathbf{M}, t) + f_i^{eq}(\mathbf{G}, t)}{3} \right). \quad (7)$$

Here, the local equilibrium  $f_i^{eq}$  is defined in the following way.

$$f_i^{eq}(\mathbf{x}, t) = w_i \rho(\mathbf{x}, t) \left[ 1 + \frac{\mathbf{c}_i \cdot \mathbf{u}(\mathbf{x}, t)}{c_s^2} + \frac{1}{2} \left( \frac{\mathbf{c}_i \cdot \mathbf{u}(\mathbf{x}, t)}{c_s^2} \right)^2 - \frac{|\mathbf{u}(\mathbf{x}, t)|^2}{2c_s^2} \right], \quad (8)$$

where  $w_i$  is the lattice weight as following:

$$w_i = \begin{cases} \frac{2}{36}, & i = 1, \dots, 6, \\ \frac{1}{36}, & i = 7, \dots, 18, \\ \frac{12}{36}, & i = 19, \dots \end{cases}$$

In this paper, the value of lattice speed of sound is  $c_s = 1/\sqrt{3}$  for the triangular mesh. Density  $\rho(\mathbf{x}, t)$  and velocity  $\mathbf{u}(\mathbf{x}, t)$  are calculated in the following way using  $f_i(\mathbf{x}, t)$  and  $\mathbf{c}_i$ .

$$\rho(\mathbf{x}, t) = \sum_{i=1}^{19} f_i(\mathbf{x}, t), \quad (9)$$

$$\mathbf{u}(\mathbf{x}, t) = \frac{1}{\rho(\mathbf{x}, t)} \sum_{i=1}^{19} f_i(\mathbf{x}, t) \mathbf{c}_i. \quad (10)$$

Substituting Eqs. (3), (4) and (7) to Eq. (2), we get

$$f_i(\mathbf{P}, t + \Delta t) = f_i(\mathbf{P}, t) + \frac{\Delta t}{A_p} (\text{Collision}_p - \text{Flux}_p) \quad (11)$$

which is the value of  $f_i$  on  $\mathbf{P}$  for the next time step. Here  $A_p$  is the total area,  $\text{Collision}_p$  is the sum of Eq. (7), and  $\text{Flux}_p$  is the sum of Eq. (4), respectively. In other words, let  $f_i(\mathbf{x}, n\Delta t)$  be denoted by  $f_i^n$ ,  $f_i^{n+1}$  is obtained. Eqs. (9) and (10) allow us to calculate  $\mathbf{u}^{n+1}$ . We add a correction step for velocity  $\mathbf{u}^{n+1}$  to eliminate outward normal vector components of  $\mathbf{u}^{n+1}$ . The outward unit normal vector  $\mathbf{n}^v$  can be calculated as

$$\mathbf{n}^v = \frac{\sum_{q \in I(\mathbf{P})} w_q \mathbf{n}_q}{|\sum_{q \in I(\mathbf{P})} w_q \mathbf{n}_q|}, \quad (12)$$

where  $I(\mathbf{P})$  is the indexed set of one-ring neighboring points of  $\mathbf{P}$ ,  $\mathbf{n}_q$  is the outward directional unit normal vector of neighboring triangle  $T_q$  and  $w_q$  is the square inverse of distance between  $\mathbf{P}$  and centroid  $\mathbf{G}$ , i.e.,  $w_q = 1/|\overline{PG}|^2$ . See [34,38,39] for further information about velocity correction. The velocity field is corrected by vector projection, Eq. (13). Fig. 3(c) shows the schematic illustration of velocity correction.

$$\mathbf{u}^{n+1} = \mathbf{u}^{n+1} - (\mathbf{u}^{n+1} \cdot \mathbf{n}^v) \mathbf{n}^v. \quad (13)$$

Finally, we update  $f_i^{eq,n}$  to  $f_i^{eq,n+1}$  using Eq. (8) with the updated density and corrected velocity field. Repeat this entire section for the next time step.

### 3. Numerical results

In this section, various experiments of fluid flows on curved 3D surfaces are demonstrated to illustrate the proposed method. Two layers of vortices are placed on a sphere to demonstrate Kelvin–Helmholtz instability (KHI) and storm behavior. Numerical experiments are conducted using various relaxation values, and the computational results are illustrated accordingly.

#### 3.1. Kelvin–Helmholtz instability on discretized coastlines

Now, we consider the KHI which consists of two layers of vortices on a sphere with a radius of  $R = 10$ . The marker function is defined as follows:

$$\phi_1(\mathbf{x}) = \begin{cases} \tanh\left(\frac{R\theta - 12.5\pi/3 - \cos(10\psi)}{0.25\sqrt{2}}\right) & \text{if } z > 0, \\ \tanh\left(\frac{-R\theta + 17.5\pi/3 + \cos(10\psi)}{0.25\sqrt{2}}\right) & \text{otherwise,} \end{cases} \quad (14)$$

where  $\theta = \cos^{-1}(z/r)$ ,  $\psi = \tanh^{-1}(y/x)$ , and  $r = \sqrt{x^2 + y^2 + z^2}$ . For the two layers of vortices, the initial velocity field is defined as follows:

$$\mathbf{u}_1(\mathbf{x}, 0) = (0.01y\phi_1, -0.01x\phi_1, 0). \quad (15)$$

We use  $\Delta t = 0.01$  to conduct simulations until  $t = 10$ . We use  $\tau_{land} = 1$  and  $\tau_{sea} = 1, 10$  to confirm the effect of the variable relaxation time value for the KHI on a sphere. Figs. 4(a) and (b) show snapshots of KHI on a sphere for  $\tau_{land} = \tau_{sea} = 1$  and  $\tau_{land} = 1, \tau_{sea} = 10$ , at time  $t = 0, 6, 10$  from left to right. The relation between the relaxation parameter  $\tau$  and the kinematic viscosity  $\nu$  in NS equation is given as  $\nu = c_s^2(\tau - 1/2)\delta_t$ , where  $c_s$  is speed of sound. The kinematic viscosity is the ratio of dynamic viscosity to the density of the fluid defined as  $\nu = \eta/\rho$ , where  $\eta$  is dynamic viscosity and  $\rho$  is density. Fig. 4 shows temporal evolution of maximum velocity  $\max\|\mathbf{u}\|$  for  $\tau_{land} = \tau_{sea} = 1$  and  $\tau_{land} = 1, \tau_{sea} = 10$ . From the result of Fig. 4(c), we can observe a gradual decrease in the velocity field overall after the occurrence of KHI within the same relaxation time  $\tau_{land} = \tau_{sea} = 1$ . On the other hand, when  $\tau_{land} = 1$  and  $\tau_{sea} = 10$ , localized turbulence occurs in areas with relatively high kinematic viscosity, such as the sea.

#### 3.2. Storm behavior at $\tau_{land} = \tau_{sea}$

Storm behaviors and evolution in different phases are popular subjects of study. Wang et al. [40] provided a dynamic model consisting of a gaseous and particle phases to describe flow in a dust storm. Bousoo et al. [41] also proposed a model for mixed flows in stormwater systems, using multiple phases such as air and water. Hession et al. [42] presented a two-phase Monte Carlo method to understand the evolution and propagation of hydrological and water quality models. Guardado-France et al. [43] studied storms from andesite sea cliffs on Isla San Luis Gonzaga, considering numerous phases depending on the geological location.

Using the same triangular mesh and coastline data, we added a storm to our globe model with a radius of 10. The center of the storm is located at  $(-2.7063, -9.0583, 3.2593)$  and the storm has a radius of 4. The two layers of vortices introduced in Section 3.1 are substituted for the numerical demonstration of storm behaviors. Fig. 5 illustrates the initial state of the globe model.

In this section, we test the storm model with the same relaxation time value  $\tau$  on all points. i.e.,  $\tau = \tau_{land} = \tau_{sea}$  and we do not distinguish whether a point is on land or sea. We calculated the results for  $\tau = 1$  and 10 until the final time  $T = 10$ . Fig. 6(a) and (b) shows the snapshots of the numerical solution on a sphere at times  $t = 0.5, 5$ , and 10. Fig. 6(c) shows that the maximum of the  $\|\mathbf{u}\|$ . We observe that at  $\tau = 10$ , compared to  $\tau = 1$ , kinematic viscosity increases, leading to accelerated fluid flow, which causes velocities to cancel out and merge, resulting in relatively large oscillations. Additionally, when  $\tau = 1$ , lower kinematic viscosity compared to  $\tau = 10$  causes velocities to monotonic decrease over time due to viscosity effects, particularly after a certain point in time.

#### 3.3. Storm behavior at $\tau_{land} \neq \tau_{sea}$

In this section, we test the storm model with land and sea having different  $\tau$  values. We fix  $\tau_{land} = 1$  and change  $\tau_{sea}$  values to 1 and 10.



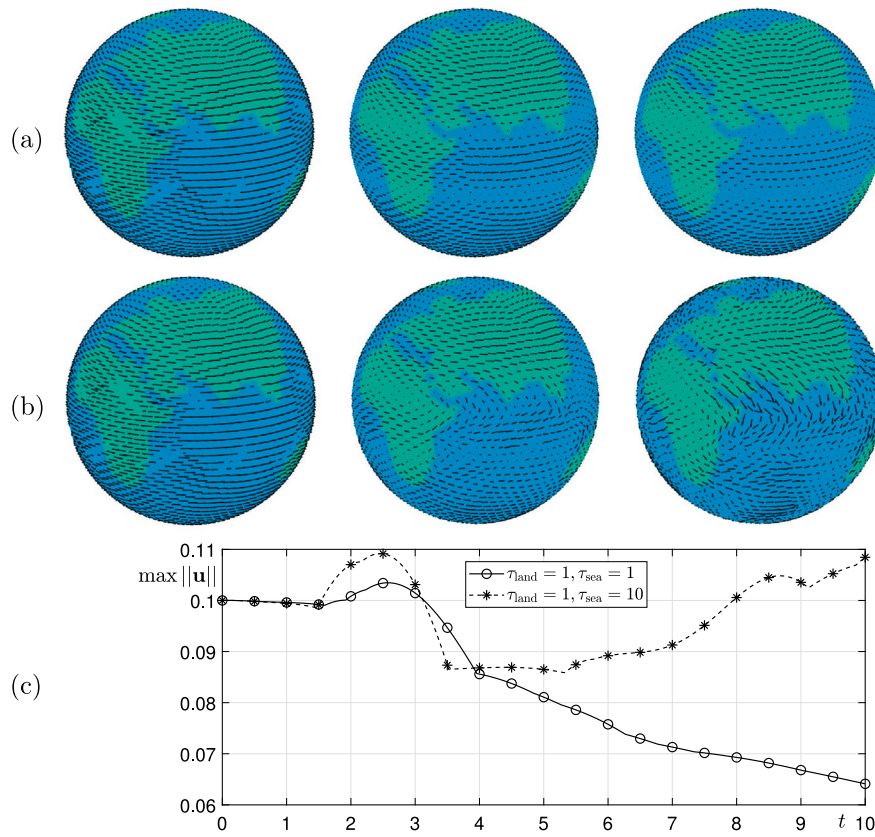


Fig. 4. Snapshots of KHI on a sphere for (a)  $\tau_{land} = \tau_{sea} = 1$  and (b)  $\tau_{land} = 1, \tau_{sea} = 10$ , at time  $t = 0, 6, 10$  from left to right. (c)  $\max ||\mathbf{u}||$ .

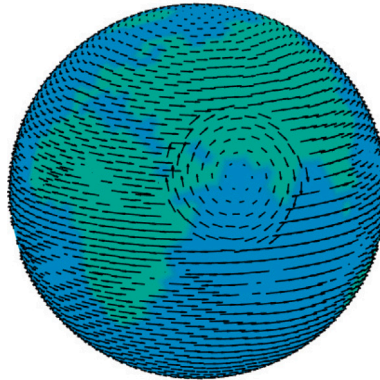


Fig. 5. Initial state after adding a storm.

Fig. 7(a) and (b) shows the snapshots of the numerical solution on a sphere at times  $t = 0.5, 5$ , and  $10$ . Fig. 7(c) shows that the maximum of the  $||\mathbf{u}||$ . We can observe that the numerical simulation results are similar to those in Fig. 4. However, the width of the oscillations is larger due to the influence of the storm.

Finally, we consider a storm landing at the southwest of North America. The two layers of vortices are not included for this experiment. The initial state is shown in the left column of Fig. 8. The parameters are unchanged as in previous simulations. In the first simulation, we set  $\tau_{land} = \tau_{sea} = 1$  and display the results in the top row of Fig. 8. The bottom row shows the results with respect to  $\tau_{land} = 3.5$  and  $\tau_{sea} = 1$ . With the increase in  $\tau_{land}$ , we observe that the velocity field dissipates faster on land.

#### 4. Conclusions

In this paper, we presented an efficient numerical method for simulating incompressible fluid flow with variable viscosity on a spherical surface. We triangulated the spherical surface and simplified the coastlines using an interpolation. Therefore, it reduced calculation and achieved high efficiency. Then, the finite volume LBM was applied to solve incompressible fluid flows. The velocity field remains tangential to the surface through the velocity correction technique at each time step in the numerical method. Validation of the proposed method was conducted through various numerical tests, such as the evolution of a two-layer vortex and a storm. Computational results showed that our method is efficient for simulating incompressible fluid flow with

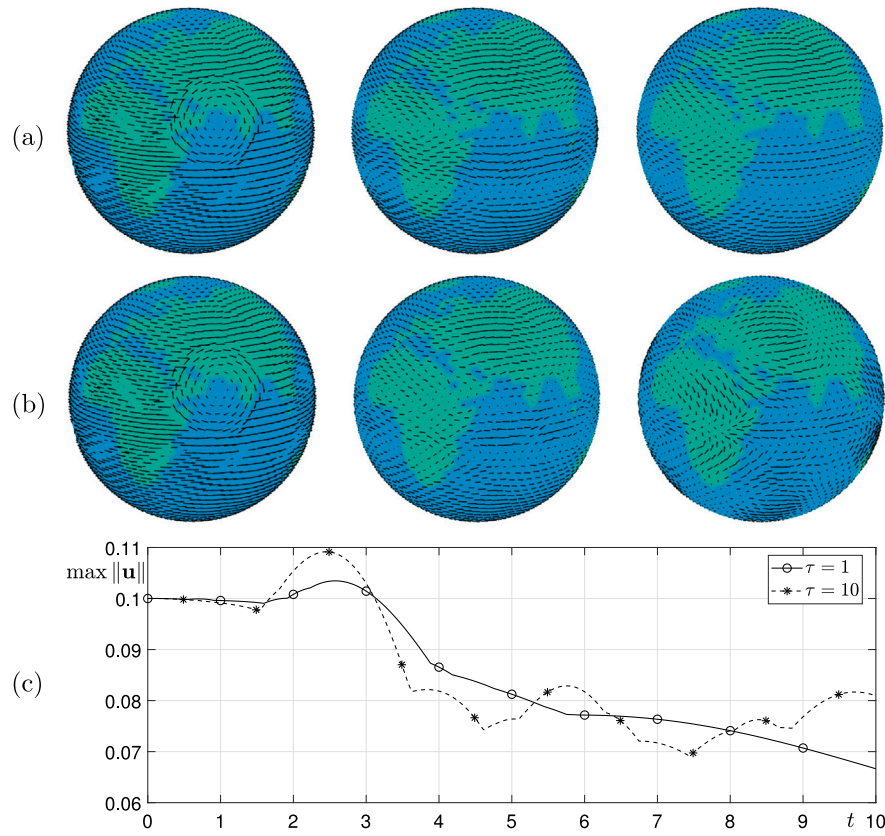


Fig. 6. Snapshots of the numerical solution on a sphere at times  $t = 0.5, 5$ , and  $10$  from left to right. (a)  $\tau = 1$  and (b)  $\tau = 10$  on both land and sea areas. (c)  $\max ||\mathbf{u}||$ .

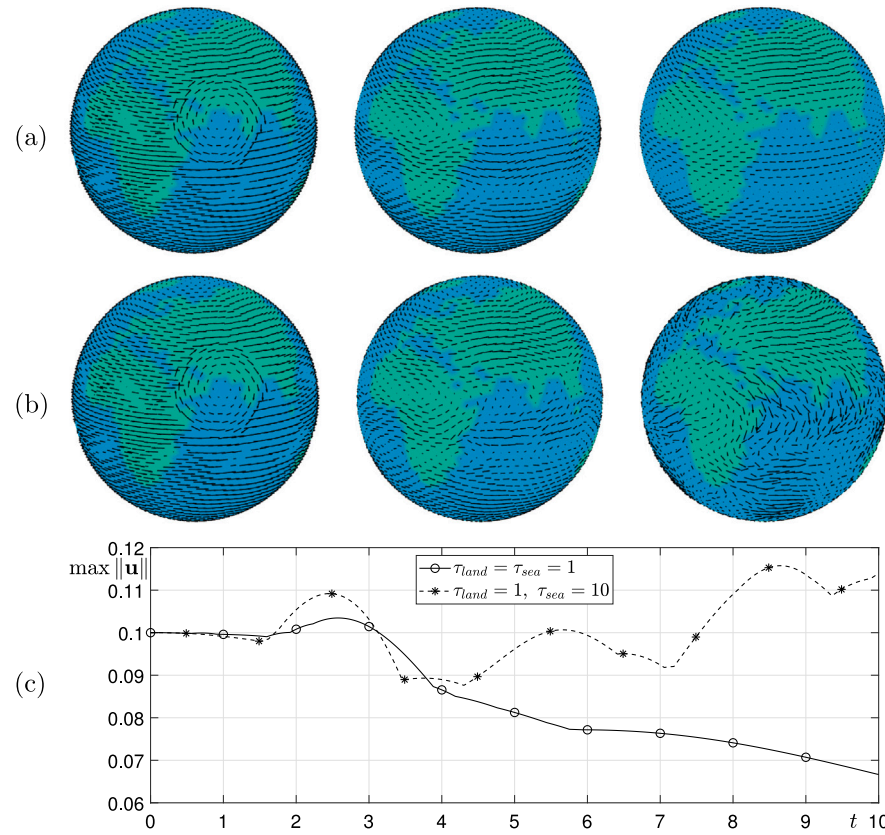


Fig. 7. Snapshots of the numerical solution on a sphere at times  $t = 0.5, 5$ , and  $10$  from left to right. (a)  $\tau_{land} = \tau_{sea} = 1$  and (b)  $\tau_{land} = 1, \tau_{sea} = 10$ . (c)  $\max ||\mathbf{u}||$ .

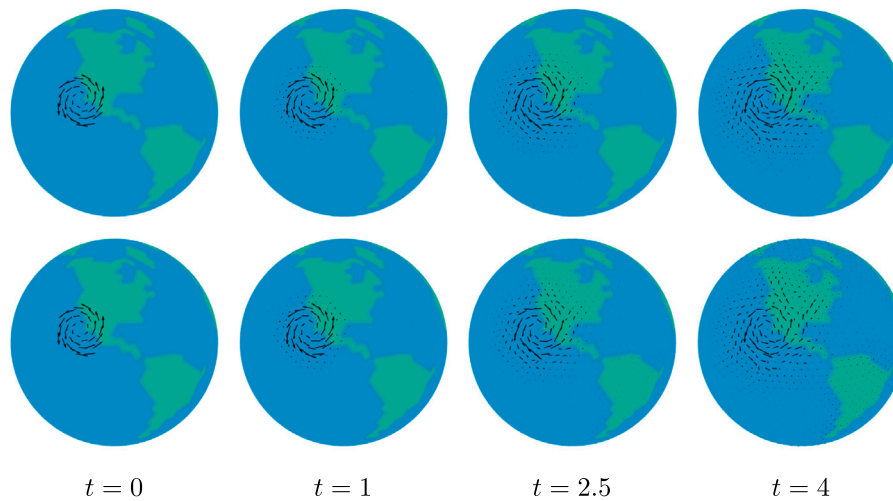


Fig. 8. Effect of  $\tau_{land}$  on a storm landing at the southwest of North America. The top and bottom rows show the results with  $\tau_{land} = 3.5$  and 1, respectively. Here,  $\tau_{sea} = 1$  is fixed.

variable viscosity. For future work, we will extend the current method to include multiphase fluid flows [44–47] and data assimilation for weather modeling [48].

#### CRediT authorship contribution statement

**Junxiang Yang:** Writing – review & editing, Writing – original draft, Visualization, Validation, Software, Methodology, Investigation, Data curation. **Seungyeon Kang:** Writing – review & editing, Writing – original draft, Visualization, Software, Resources, Methodology, Investigation, Data curation. **Youngjin Hwang:** Writing – review & editing, Writing – original draft, Resources, Investigation, Data curation. **Soobin Kwak:** Writing – review & editing, Writing – original draft, Resources, Methodology, Data curation. **Seokjun Ham:** Writing – review & editing, Writing – original draft, Software, Resources, Investigation. **Junseok Kim:** Writing – review & editing, Writing – original draft, Visualization, Validation, Supervision, Software, Project administration, Methodology, Investigation, Funding acquisition, Formal analysis, Conceptualization.

#### Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

#### Data availability

No data was used for the research described in the article.

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